New approaches to exploration for IOCG-style mineralisation, Middleback Ranges, S.A.

Thesis submitted in accordance with the requirements of the University of Adelaide for an Honours Degree in Geology.

Holly Feltus
November 2013



RUNNING TITLE

REE distributions: a new IOCG exploration tool.

ABSTRACT

Iron oxide copper gold (IOCG) systems display well-developed spatial zonation with respect to alteration assemblages, mineralogy and the distribution of rare earth elements (REE). The Middleback Ranges, South Australia, located in the Olympic Province, Gawler Craton, hosts anomalous Fe-oxide-bearing Cu-Au mineralisation, and are considered potentially prosperous for larger IOCG-style deposits. This study investigates whether the distribution of REE and other trace elements within selected minerals represents a potential exploration tool in the area. Iron-oxides (hematite and magnetite), potassium feldspar, albite and accessory minerals have been analysed by laser-ablation inductively-coupled plasma mass spectrometry (LA-ICP-MS) from two prospects (Moola and Princess) and in samples of the Myola Volcanics. The resultant multi-element datasets are compared to other IOCG systems.

The results support the presence of sizeable and/or multiple IOCG alteration envelopes within the Middleback Ranges. Significant evolving hydrothermal events resulted in hydrolithic alteration and remobilisation of REE within the Moola Prospect and Myola Volcanics.

Replacement of early magnetite by hematite (martitisation) in the Myola Volcanics is accompanied by an influx of REE visible on LA-ICP-MS element maps showing partial martitisation at the grain-scale. It is thus inferred the initial generation of magnetite must have pre-dated introduction of oxidised, REE-enriched hydrothermal fluids into the system. Sulphide assemblages observed within the Moola Prospect are complex and record sequential recrystallisation under evolving fS_2 and fO_2 conditions. Trace minerals, cycles of brecciation and replacement, and distributions of REE within minerals are similar to that observed in other IOCG domains. The Princess Prospect displays REE distributions in minerals which are dissimilar to the Moola Prospect, the Myola Volcanics and also those reported from other IOCG domains. This is interpreted as indicating that the Moola Prospect and Myola Volcanics in the south of the Middleback Ranges are more prospective IOCG targets.

KEYWORDS

Middleback Ranges, Iron-Oxide Cu-Au (IOCG), Rare Earth Elements (REE), incompatible elements, exploration, alteration.

REE distributions: a new IOCG exploration tool.

TABLE OF CONTENTS

List of Figures
List of Tables
Introduction
Background
IOCG mineral systems
Regional Geology12
IOCG mineralisation in the Middleback Ranges
Methods
Results
Discussion
Comparison of the Princess and Moola Prospects and Myola Volcanics 62
REY distributions and their petro genetic and exploration significance 63
Towards a preliminary genetic model
Conclusions
Acknowledgments
References
Appendix

LIST OF FIGURES

Figure 1 Regional-scale map showing the Middleback Ranges within the Olympic
Province. Major IOCG deposits are located
Figure 2 Schematic cross-section illustrating idealised alteration zonation in IOCG
deposits (Hitzman et al. 1992)
Figure 3 Geological sketch map of the Middleback Ranges, also showing the three
sampling localities
Figure 4 (a-h) Photographs of hand-specimens of typical lithologies (scale-bars: 1 cm).
(a) Myola Volcanics porphyry rhyolite (sample: MV01); (b) Felsic breccia (sample:
ML01); (c) Banded felsic gneiss; granitic veinlet is seen on the RHS of image (sample:
ML02); (d) Granite veinlet (sample: ML06); (e) Granite (sample: ML07); (f) Flow
banded rhyolite-dacite volcaniclastic (sample: ML15); (g) Hematite breccia (sample:
PS01); (h) Metasedimentary rock (sample: PS03)
Figure 5 (a-h) Petrographic images of typical lithologies; all transmitted cross polarised
light images except g) (reflected light cross polars). (a) Myola Volcanics porphyry
rhyolite (sample: MV01), Microcline porphyryblast within a fine grained matrix
comprised of quartz, feldspar and sericite. Chlorite, sericite and rutile are associated
with feldspars. Fabric is defined by elongated zones of coarse-grain minerals and
martite stringers (opaque mineral). (b) Felsic breccia (sample: ML01); Chlorite
dominated breccia infill. RHS of image is a clast of granite composed of quartz, altered
feldspars and minor zircon. (c) Banded felsic gneiss (sample: ML02); Fabric is defined
by coarse-grain chlorite and sericite. Rock is dominated by quartz and altered feldspars.
(d) Granite veinlet (sample: ML06); Central mineral is heavily chlorite altered
hornblende surrounded by quartz, muscovite and altered feldspars. (e) Granite (sample:
ML07); Granite is dominated by quartz, heavily altered microcline, plagioclase,
chlorite, sericite and titaniferous hematite. Opaque mineral in centre of image is
chalcopyrite. (f) Flow banded rhyolite-dacite volcaniclastic (sample: ML15); Fine
grained volcaniclastic is dominated by quartz, feldspar, muscovite, titaniferous hematite
and ilmenite. (g) Hematite breccia (sample: PS01); Bladed hematite (fine- and coarse-
grained in infill and clasts respectively) dominates the breccia. (h) Metasedimentary
rock (sample: PS03); The very fine-grained rock is dominated by quartz, chlorite and
sericite with minor kutnohorite and quartz-carbonate crackle veins. Scale bar: 500 μm.
Figure 6 (a-k) Back-scatter electron images showing accessory minerals. a)
Equigranular rutile (Ru) associated with monazite (Mon) from the Princess Prospect
metasediments; rutile is potentially hydrothermal (sample ID: PS06). b) Zoned and
broken zircon (Zrc) is common throughout the sample suite (sample ID: ML04 - felsic
banded gneiss). c) The felsic volcaniclastic rock has symplectic rutile (darker grey) and
hematite (Hm) (brighter grey) after ilmenite (Il). Bladed Ti-poor hematite is also
observed in this image (brightest mineral) (sample ID: ML14). (d) Myola Volcanics –
late-stage pyrolusite (Pyl); zonation is apparent in image however analysis showed no
obvious chemical variation (sample ID: MV01). e) & f)) Granitic veinlet - hematite with
ilmenite exsolution lamellae is associated with a homogenous course grained rutile.
Inset f) is a close up of exsolution textures (sample ID: ML03). g) & h) The contact
between felsic banded gneiss and granitic veinlet commonly has rutile and ilmenite;
inset h) Rutile and ilmenite close up (sample ID: ML04). i) This rutile from the Princess
Prospect metasediments may represent a corroded detrital grain (sample ID: PS06). j)

REE distributions: a new IOCG exploration tool.

Rutile in the Myola Volcanics is often associated with titanite (Ttn) and Fe-oxides
(sample ID: MV01). k) The rutile in the felsic volcaniclastic rock is observed to replace
Fe-oxides (sample ID: ML17).
Figure 7 (a-f) Petrographic images – (a, b & e) reflected light, plane polars light; (c, d &
f) transmitted light, cross polars light. All images are from the Moola Prospect. a) This
image from the hybrid zone shows pyrite (Py) and marcasite (Ma) with late-stage
chalcopyrite (Cpy) (sample ID: ML15). b) Corroded pyrite with late-stage chalcopyrite
was observed within a quartz-carbonate vein in the granitic veinlet; fine-grain sphalerite
was also observed within the vein (sample ID: ML03). c) Late-stage carbonate vein
within the granitic veinlet (sample ID: ML04). d) Hornblende (Hbl) is commonly
altered to chlorite (Chl); feldspars are often altered by chlorite and sericite. Muscovite
(Mu) can be both coarse- and fine-grained (sample ID: ML04). e) Titaniferous hematite
(Hm) showing ilmenite exsollution lamellae (<1 μm across and ~10 μm long) (sample
ID: ML01). f) Simple twinning of microcline (Ksp) with sericite (Mu) alteration within
granitic clast in felsic breccia (sample ID: ML01)
Figure 8 (a-i) Back-scatter electron images of Fe-oxides and sulphides; all images are
from within the silicic breccia at the contact between the granite and felsic
volcaniclastic in the Moola Prospect except for c), f), g) & h). a) Early pyrite (Py) has
chalcocite (Cc) along fractures. Massive chalcocite is associated with wittichenite (Wit)
and has late-stage delafossite (Dlf) (sample ID: ML13). b) Kutnohorite (Ku) is zoned
with respect to Mn; coeval growth is supported by Cu-sulphate (?) inclusions within the
core of the kutnohorite grain. (sample ID: ML13). c) Pyrite within the Princess
Prospect metasedimentary rocks is zoned with respect to As (sample ID: PS12). d)
Early pyrite is proximal to minor sphalerite (Sp); covellite (Co) is observed growing
into the void, and late-stage Kutnohorite and Cu-sulphates are forming within voids.
Relationships in b) and d) may support a relationship between late-stage remobilisation
of Cu and Mn (sample ID: ML13). e) Chalcocite replacing pyrite (sample ID: ML13). f)
Hematite breccia: Princess Prospect. Bladed hematite (Hm) is rimmed by a late-stage
magnetite (Mt) that has elevated REY in comparison to the Hm (sample ID: PS02). g)
Moola Prospect: granite. Late-stage uraninite (U) and monazite (Mon) is commonly
closely associated with chalcopyrite and apatite (Ap) and is proximal to kutnohorite
veins (sample ID: ML10). h) Moola Prospect: granite veinlet. Corroded pyrite is
rimmed by late-stage uraninite (sample ID: ML03). i) Native dendtritic copper (sample
ID: ML13)
Figure 9 (a & b) Reflected light images in plane polarised light; martite texture –
hematite (Hm) replacing magnetite (Mt). This feature is predominant in the Myola
Volcanics. c) Chondrite-normalised REY fractionation trends for fresh magnetite,
slightly altered magnetite and martite show the degree of martitisation is associated with
REE enrichment
Figure 10 Chondrite-normalised REY fractionation trends for titaniferous hematite
(Moola Prospect). Note irregular distribution; there is a slight increase of REY down-
hole
Figure 11. Back-scatter electron image showing the relationship between titaniferous
hematite, rutile and ilmenite in the flow-banded rhyolite found towards the base of
ML001DD. Primary titaniferous hematite is observed throughout the sample; this coarse
grained rutile shows fracturing and replacement by late-stage impure ilmenite.
Chondrite-normalised REY fractionation trends for a) ilmenite, b) rutile and c) hematite,
clearly indicating that the late-stage ilmenite is associated with an influx of REY 35
cicarry murcarrig that the late-stage inhelite is associated with all limux of RE 1 53

REE distributions: a new IOCG exploration tool.

Figure 12 Chondrite-normalised fractionation trends for a) albite and b) microcline. See
text for additional information
Figure 13 Chondrite-normalised REY fractionation trends for accessory minerals. a)
Apatite, containing significant amounts of REY. Granitic apatite within the Moola
Prospect displays a concave trend similar to Hillside ore-stage altered skarn (Ismail et
al. in press). b) Titanite is particularly HREE-rich, a feature unique to a subset of
environments within IOCG and skarn systems. c) & f) Rutile, showing variable REY
plots throughout the sample suite; see text for explanation. e) Late-stage pyrolusite is a
significant carrier of REY and is LREE-enriched. e) Kutnohorite REY patterns differ
between samples. Note: standard used for kutnohorite analysis did not contain Tb, Y,
Tm or Yb – these elements are not displayed
Figure 14 a) & b) Rb-Sr-Ba ternary plots for feldspars. Note albite is relatively rich in
Sr in comparison to microcline. c) Rb-Ba binary plot showing distinct trends for
microcline from felsic volcaniclastics, rhyolite porphyry, and granite. These plots allow
for discrimination among lithologies. d) & e) U-Th-Pb ternary plots for feldspars,
showing that albite has increased concentrations of U and Th with respect to Pb 55
Figure 15 a) Reflected light image displaying the martite texture of hematite (Hm)
replacement of magnetite (Mt). Remaining images are LA-ICP-MS element maps for
the martite grain shown in (a). The degree of martitisation correlates with REE
enrichment (particularly LREE). A moderate correlation can be seen between
martitisation and the concentrations of Mn and Zn. Maps showing the distributions of V
and Co illustrate their presence in Fe-oxides. Scales are in counts per second
(logarithmic scale). 57
Figure 16 a) Back-scatter electron image of kutnohorite grain displaying compositional
zonation (scale bar 1 mm). Remaining images are LA-ICP-MS element maps of this
grain. Mn, Mg, Fe and Ca maps show that the grain-scale compositional zonation is
largely attributable to major variations in Mn content. Kutnohorite is also zoned with
respect to, and is a significant carrier of various metals and incompatible elements.
Scales are in counts per second (logarithmic scale)
LIST OF TABLES
Table 1 Archaean to Paleoproterozoic stratigraphy of north-eastern Eyre Peninsula
(Parker 1993)
Table 2 Petrographic summary of main lithologies
Table 3 Summary of LA-ICP-MS trace element data for Fe-oxides (ppm)
Table 4 Summary of LA-ICP-MS trace element data for feldspar (ppm)
Table 5 Summary of LA-ICP-MS trace element data for rutile (ppm)
Table 6 Summary of LA-ICP-MS trace element data for apatite (ppm)
Table 7 Summary of LA-ICP-MS trace element data for kutnohorite (Ku), pyrolusite
(Pyl) and titanite (Ti) (ppm)
Table 8 Results of Zr-in-rutile geothermometry using the calibration of Watson et al.
(2006)
Table 9 Electron probe microanalytical data for chlorite. Estimated formation
temperature is calculated based on the calibrations of Cathelineau (1988) and Jowett
(1991)
(±>>±/