



# LANDSCAPE EVOLUTION OF THE UMBUM CREEK CATCHMENT, WESTERN LAKE EYRE, CENTRAL AUSTRALIA.

#### VICTOR G. WACLAWIK

**BSc. (Hons) (FLINDERS)** 

Thesis submitted as fulfillment of the requirements for the degree of Doctorate of Philosophy

Discipline of Geology and Geophysics School of Earth and Environmental Science University of Adelaide Australia

## **TABLE OF CONTENTS**

Αb	strac	ct	i
De	clara	ationation	ii
Acknowledgements			iii
0			
1.	Intr	roduction	1_1
١.		Rationale	
		Aims	
		Objectives	
		Background	
		Study Area	
		Stratigraphy	
		Quaternary Climate	
		Structure and Tectonics	
	1.9	Mound Springs	1-19
		O Dune Orientations	
	1.1	1 Geomorphology	1-28
2.		thodologythodology	
	2.1	Seismicity in the Lake Eyre Region	
		2.1.1 Datasets	
		2.1.2 Statistical Methods	
	2.2	Remote Sensing	
		2.2.1 Landsat 5TM	
		2.2.2 Analysis of Multispectral ASTER Data	
		2.2.3 JERS-1 SAR	
	2.2	2.2.4 JERS-1 SAR and ASTER Image Melding	
	2.3	Tectonic Influences on Geomorphology	
		<ul><li>2.3.1 Digital Elevation Model and Longitudinal Stream Profiles</li><li>2.3.2 Morphometric Analysis</li></ul>	
		2.3.3 Differential GPS Survey	
		2.3.4 Geological Cross-sections	
	24	Geological Structure Mapping	
	۷. ۱	2.4.1 Standard Geological Methods	
		2.4.2 Ground Penetrating Radar	
	2.5	Regolith Geology	
		2.5.1 Regolith Mapping	
	2.6	Summary	
3.	Sei	ismicity in the Lake Eyre Region	3-1
		Introduction	
	3.2	Tectonic Model	3-1
	3.3	Results	3-1
	3.4	Discussion	
		3.4.1 Hydroseismicity	
		Further Implications	
	3.6	Conclusions	3-12
4.	Rei	mote Sensing	
	4.1	0.401.0 2 4.4	
		Magnetic Data	
		Radiometric Data	
		Basement Structural Interpretation	
	4.5	Landsat 5TM	
		4.5.1 False Colour Composites	4-5

	4.5.2 Band Ratios	4-10
	4.5.3 Landsat Summary	4-16
	4.6 ASTER Image Analysis	
	4.6.1 Hawker Springs	4-19
	4.6.2 Neales Fan	
	4.6.3 Sunny Creek	4-36
	4.6.4 ASTER Summary	
	4.7 JERS-1 SAR	
	4.8 JERS-1 SAR and ASTER Image Melding	
	4.9 Remote Sensing Summary	
5.	Tectonic Influences on Geomorphology	5-1
	5.1 Drainage Analysis	5-1
	5.1.1 Neales River and Brown Creek	5-1
	5.1.2 Umbum Creek and Sunny Creek	5-4
	5.1.3 Davenport Creek	
	5.1.4 Hawker Creek	
	5.2 River Channel Morphology	5-4
	5.3 Longitudinal Profiles	
	5.4 Slope and Aspect	
	5.5 Drainage Density	
	5.6 Hypsometric Analysis	
	5.7 Elevation Histogram	
	5.8 Geomorphometry	
	5.9 Rangefront Sinuosity	
	5.10 Flooding Surface Analysis	
	5.11 Differential GPS Survey	
	5.11.1 Neales River Cross-channel Profiles	
	5.11.2 Umbum Creek Cross-channel Profiles	
	5.11.3 Neales Fan Axial Profiles	
	5.11.4 Neales Fan Cross-profiles	
	5.12 Geological Cross-sections	
	5.12.1 Idealised Geological Section Neales Fan Looking North	5-30
	5.12.2 Idealised Geological Section Neales Fan Looking West	5-30
	5.13 Geomorphological Summary	5-30
6.	Geological Structure Mapping	6-1
	6.1 Neales Cutbank	
	6.1.1 Lineaments on the Neales Fan	6-7
	6.1.2 Survey Results	6-11
	6.1.3 Climatic Control	6-11
	6.1.4 Structural Control	6-11
	6.1.5 Lithological Control	6-15
	6.2 Model Evaluation	
	6.3 Umbum Creek	
	6.4 Geological Structure Summary	
	6.5 Ground Penetrating Radar (GPR)	
	6.5.1 GPR Summary	
	oion oi n ou maight ann an	6 20
7.	Regolith Geology	7-1
٠.	7.1 Regolith Mapping	
	7.2 Landforms	
	rizir ina zavanpari nangaa mangarat zana ana maani mangarat zasarpin	
	7.2.2 Alluvial Fans	
	7.2.3 Reworking of Jurassic Sediment	
	7.2.4 Tilted Silcrete Outcrops	
	7.2.5 Silcrete and Well-rounded Cobble Float	
	7.2.6 Mound Springs	7-17

		7.2.7	Alluvial Overbank Plains and Channel Cross-section Morphology	7-19
		7.2.8	Claypans	
		7.2.9	Umbum Overflow	
		7.2.10	Four Hills and Surrounds	
		7.2.11	Sunny Creek/Douglas Creek Interfluve	
		7.2.12	Midla Milda and Wadlalgla Hills	
		7.2.13	Dunes	
		7.2.14	Neales Fan Drainage	
		7.2.15	Neales Fan Palaeochannels	
		7.2.16	Deltas/Terminal Splays	
	7.0	7.2.17	Carbonate-cemented Conglomerate Palaeochannel Equivalents	
	1.3		Comparison	
		7.3.1	Sand Dunes and Gravity	
		7.3.2	Regolith and Magnetics	
	7.4	7.3.3 Regolith	Regolith Landforms and Magnetic Lineaments	
		J	•	
8.			e Processes	
			ction	
	8.2	Marine :	Sedimentation, Iron Di-sulphide Formation and Acid Sulphate Soils	
		8.2.1	The Global Biogeochemical Carbon & Sulphur Cycle	
		8.2.2	Acid Sulphate Soils and Silcrete	
			Processes	
	8.4		Plains	
		8.4.1	Wind Erosion	
		8.4.2	Physico-chemical Weathering	
		8.4.3	Colluvial Wash and Banded Vegetation	
		8.4.4	Upward Migration of Particles	
	٥.	8.4.5	Accretionary Mantles	
	8.5		Plain Sampling	
		8.5.1	JH002 & JH002a	
		8.5.2	JH005, JH006 & JH007	
		8.5.3	JH008 & JH009	
		8.5.4	JH010, JH011 & JH012	
		8.5.5 8.5.6	Overprinting of Processes Influencing Gibber Plain Formation	8-17
	0.4		Sieve Analysis: Pit JH005	
	0.0	8.6.1	n and Gypcrete	
	0 7		Modes of Formation of Gypsum in the Umbum Creek Catchmente Processes	
	0.7	8 7 1	Dryland Catchments and Ephemeral Streams	
		8.7.2	Headward Erosion and Gully Formation	
	0 0		ape Processes Summary	
			ion	
	0.7	Discuss		
9.	Lar	ndscape	e Evolution Model	9-1
	9.1	Landsca	ape Evolution	9-1
		9.1.1	Pre-Cainozoic	
		9.1.2	Cainozoic	9-2
		9.1.3	Landscape Evolution of Umbum Creek Catchment	9-4
		9.1.4	Landscape Evolution of the Neales Fan	9-8
10	Dis	CUSSINI	n & Conclusions	1∩₋1
10.	10.1 Discussion			
			ions	
	10.2	10.2.1	Landscape Evolution	
		10.2.1	Neotectonic Activity	
	Б.:			
11.	RID	поgrap	hy	11-1

Appendices	
Appendix A: Mineral Properties	A-1
Appendix B: Stream Longitudinal Profiles	
Appendix C: Regolith Geology of the Umbum Creek Catchment	
Appendix D: Data Locations	
•••	

## **LIST OF FIGURES**

Chapter 1: Introduction Figure 1.1: Landsat 7 mosaic of the Australian continent (courtesy Geoscience Australia) indicating major structural lineaments passing through the study area that broadly align to continent-wide structural corridors (Campbell & O'Driscoll, 1989)
Figure 1.2: Location diagram showing the position of major features within the study area on the western fringe of Lake Eyre1-3
Figure 1.3: Location Map showing the extent of the Lake Eyre Basin, Lake Eyre and major rivers (from Kotwicki and Allan, 1998)1-6
Figure 1.4: Map of Australian climatic zones showing the influence of the northern tropical monsoon system and the mid-latitude westerly circulation system across Australia. (From Nanson and Price, 1998)1-7
Figure 1.5: Spring location diagram showing springs clustered in the upper section of the catchment1-8
Figure 1.6: Diagram of Quaternary climatic change in Australia showing the fluctuation of climatic zones across the continent (From Williams, 1984)1-12
Figure 1.7: Location of the contempoary monsoonal shear zone showing the point of origin of tropical cyclones (From Croke <i>et al.</i> , 1999)1-14
Figure 1.8: Diagram of Australian tectonic structure showing the arrangement of major structural features of the Australian continent (From Gale, 1992)1-15
Figure 1.9: Diagram of Lake Eyre structural lineaments showing a strong divergence into NW- and NE- trending sets (From Firman, 1975)1-16
Figure 1.10: Interpretation of a magnetic survey showing the trends of basement features and sedimentary basins containing Mesozoic sediment (After Teluk, 1974)1-18
Figure 1.11: Schematic of Great Artesian Basin showing the development of springs at the edge of the basin as a result of the hydraulic gradient (From Thomson and Barnett, 1985)1-20
Figure 1.12: Map of artesian water quality in the Lake Eyre region showing highly corrosive waters associated with the western province (From Thomson and Barnett, 1985)1-21
Figure 1.13: Map of mound spring locations showing the relationship to springs and structural features (From Aldam and Kuang, 1989)
Figure 1.14: Diagrams illustrating the structural control of springs by faults. a) Etadurrana Springs. b) Section through the study area showing how downcutting streams intersect structural features to create springs. c) Billa Kalina Springs. d) Strangways Springs showing the tilting of fault blocks to create springs. Seismic line 86-AED shows interpreted structure beneath Dalhousie Springs (From Aldam & Kuang, 1989). 1-23
Figure 1.15: Idealised mound spring cross-section showing the association of springs with faults (From Thomson and Barnett, 1985).
Figure 1.16: Map of Australian dune trends showing the whorl-shaped distribution of trends across the continent resulting from Pleistocene wind directions (From Nanson <i>et al.</i> , 1995)1-27
Chapter 2: Methodology Figure 2.1: Location diagram of Landsat 5 TM data showing the position of the Landsat tiles used for analysis (Landsat 7 ETM mosaic image courtesy of Exxon-Mobil)
Figure 2.2: Location diagram of ASTER data showing the position of the ASTER tiles used for analysis (Landsat 7 ETM mosaic image courtesy of Exxon-Mobil)2-5
Figure 2.3: Delimitation of catchments within the study area showing the locations of catchment boundaries 2-10
Chapter 3: Seismicity in the Lake Eyre Region Figure 3.1: Interpreted basement faults in the study area showing the spatial correlation with earthquake epicentres. The number of stations that record the event and the position of these stations influence epicentre location. For the South Australian seismic network, spatial error is likely to be within a 25-50km radius. Seismic stations are located to the south and southeast of this region and thus the error is likely to plot in a north to northeasterly direction

Figure 3.2: Regional earthquake data as compared to volumetric inflows into Lake Eyre through time (1979-2002) showing little correlation between earthquake datasets and calculated proxy inflows to Lake Eyre
Figure 3.3: Regional earthquake data as compared to estimated lake levels in Lake Eyre (1979-2002) showing little correlation between earthquake datasets and calculated lake levels
Figure 3.4: Covariance plot of Frequency versus Inflow data showing a peak in correlation at a delay of six years.
Figure 3.5: Plots of Annual Frequency versus Average Annual Inflow and Average Annual Lake Level versus Annual Frequency shifted backwards by 6 years showing a stronger correlation between earthquake datasets and both lake level and proxy inflow data
Figure 3.6: Jackboot Bore location diagram showing the GAB monitored wellfields network. Jackboot Bore is circled (From Sibenaler <i>et al.</i> , 2000)
Figure 3.7: Comparison of regional hydrologic datasets with Jackboot Bore RSWL. Inflow displays a minor correlation with bore pressure, as does Lake Level
Figure 3.8: Comparison of Frequency with Jackboot Bore RSWL showing a stronger correlation when Frequency is time-shifted by 6 years
Chapter 4: Remote Sensing Figure 4.1: Gravimetric Imagery showing broad arrangement of basement structure. High gravimetric levels are associated with granitic bodies and Proterozoic basement, low gravimetric levels are associated with sedimentary basins. The basement structureappears to control the alignment and distribution of streams in the study area. 4-2
Figure 4.2: Total Magnetic Intensity image displayed as (a) greyscale image with sun-shading. (b) Interpreted basement structural lineaments identified and displayed as red (NW-SE) or white (NE-SW) lines. Streams are shown in blue4-3
Figure 4.3: (a) Radiometric imagery mosaic of surveys conducted within the study area showing potassium, thorium and uranium distribution in the landscape for exploration leases investigated in the region; potassium in red, thorium in green and uranium in blue. (b) Streams displayed in white for reference
Figure 4.4: Structural interpretation of geophysical data showing the broad arrangement of basement structure. Interpreted basement structural lineaments are identified and displayed as white lines. It appears that the basement structure controls the alignment and distribution of streams
Figure 4.5: Landsat 5 TM RGB321 natural colour image displaying identifiable landforms including 1. sand dunes; 2. ranges; 3. mound springs; 4. claypans; 5. channels; 6. terminal splay complexes; and, 7. lakes 4-7
Figure 4.6: Landsat 5 TM False Colour Composite image RGB741 showing areas of bedrock highlighted by band 7 in red and clays highlighted by band 4 in green. The dark orange area in the North of the image is a plain covered by heavily desert-varnished silcrete gibber. Alluvial fans flank the base of the Davenport Ranges and are composed of quartz sandstone, except in the south where they consist of well-rounded, poorly sorted mixed clasti sediments that grade into heavily desert-varnished silcrete gibber. Mound springs and associated streams can be seen as pale patches eroding into the flanks of the alluvial fans near the centre of the image
Figure 4.7: Detail Landsat 5 TM False Colour Composite Image RGB741 (dashed box on Figure 4.6) showing distribution of clay sediments across the surface of the Neales Fan. Clay pans are absent along the northern third of the Neales Fan. Claypans populate the middle third of the Neales Fan and form two arms that may represent mud-plugs residing in palaeochannels
Figure 4.8: Landsat 5 TM False Colour Composite image RGB532 showing large areas dominated by dunes and the effects of erosion on Mesozoic sediments in pale purple-blue associated with gypsum-rich colluvium and gypsite capped escarpments
Figure 4.9: Landsat 5 TM False Colour Composite image RGB573 showing the delimitation of alluvial units. Reworked Jurassic fluvial sands are conspicuous as pale yellow bands carried by modern alluvial systems from the northern and southern ends of the Davenport Ranges4-12
Figure 4.10: Detail of a pebble lag composed almost entirely of rounded quartz clasts derived from reworked fluvial Jurassic Algebuckina Sandstone. Scale+10 cm. (0619941E 6828652N)4-13
Figure 4.11: Landsat 5 TM False Colour Composite image RGB754 showing differences in the spectral response of surface lag units probably reflecting different degrees of desert varnish coatings4-14

Figure 4.12: Landsat 5 TM band ratio image Bands 3/1 showing the presence of iron oxides in the landscape. Sand dunes show a high response due to the presence of hematite grains that impart the red colour to dunes in this region. Alluvial fans and gibber plain surfaces show a high response due to the heavy desert varnish that coats these sediments. This varnish is deep red in colour and is likely to contain iron oxides. Sand dunes on the Neales Fan are not illuminated by this ratio and are unlikely to contain iron, suggesting a different source of origin to those dunes of the surrounding region
Figure 4.13 Landsat 5 TM band ratio image Bands 7/5 showing the presence of clay-forming minerals in the landscape as bright areas. These are associated here with stream channels, claypans and depositional depressions. There is also a geological unit in the Davenport Ranges that possesses a strong clay signature
Figure 4.14: Landsat 5 TM False Colour Composite of Band Ratios RGB3/1, 5/4, 5/7 designed to highlight iron (in red) silica (in green) and aluminium hydroxides (in blue). Silica- and aluminium hydroxide-rich rocks form the ranges and dominate the streams. Dunefields are dominated by iron and silica whereas dunes on the Neales Fan are dominated by silica. Gibber plains covered by silcrete have a high silica response and appear as green. Floodouts are associated with aluminium hydroxides
Figure 4.15: Location diagram of ASTER data showing the position of the ASTER tiles used for analysis 4-20
Figure 4.16: Non-Negative Principal Component Analysis (NNPCA) of Hawker Springs ASTER scene. The colour scale on the right-hand side represents 18 enforced endmembers. The magenta endmember highlights areas of fine-grained clays and evaporites typically associated with underlying Mesozoic sediment. The orange-yellow endmember highlights alluvial fans fringing the northern end of the Davenport Ranges. Sediment extending from the southern Davenport Ranges in a northeasterly direction is represented by the green-yellow endmember. The green endmember highlights erosional plains mantled with longitudinal dunes. A band of clouds is indicated by the dashed line
Figure 4.17: Independent Component Analysis (ICA) results for Hawker Springs showing Mesozoic sediments exposed by erosion in orange and magenta; fine-grained clays and evaporites associated with mound springs, floodouts and claypans in red; degraded alluvial fan sediments in pale blue; and, relict colluvial plains in dark blue.  4-22
Figure 4.18: Rowan's Ratios for RGB muscovite/illite/smectite, calcite, alunite/pyrophyllite for Hawker Springs ASTER scene, showing a strong response in the calcite ratio from dolomitic units in the Proterozoic sediments of the Davenport Ranges. Streams have a strong response in both the calcite ratio and alunite/pyrophyllite suggesting that stream sediments are derived from the Proterozoic sediments in the stream provenance rather than contribution from the surrounding landscape. Sand dunes display a strong alunite/pyrophyllite signal 4-24
Figure 4.19: Rowan's Ratios for RGB muscovite/illite/smectite, dolomite, alunite/pyrophyllite for Hawker Springs ASTER scene, showing a strong response in the dolomite ratio from dolomitic units in the Proterozoic sediments of the Davenport Ranges. This response is further reflected in streams draining these bedrock units. Alluvial fans flanking the ranges possess a strong response in the muscovite/illite/smectite ratio with a similar response from source rocks in the Davenport Ranges. The southern sediments flanking the range and areas of gibber plain show a strong response in the dolomitic ratio. This is probably due to similar absorption features between desert varnish and dolomite spectra
Figure 4.20: Rowan's Ratios for RGB muscovite/illite/smectite, dolomite, calcite for Hawker Springs, showing a strong distinction between sources of sediment. Proterozoic sediments and alluvial fans have large amounts of muscovite-type minerals; the surrounding gibber plains appear to be dominated by a dolomite signal This is probably due to similar absorption features between desert varnish and dolomite spectra. A unit in the Proterozoic sediments of the Davenport Ranges has a strong calcite signal. This is also evident along streams and indicates that this unit is the source for much of the stream sediment
Figure 4.21: Continuum removed comparison of dolomite and desert varnish spectra showing similar absorption features
Figure 4.22: Mahalanobis Distance Classification image for ASTER Hawker Springs scene showing a strong correlation between spectral properties and landform type. There is a clear distinction between alluvial fans (in red) flanking the ranges and the sediment flanking the ranges in the southeast of the image (in orange). There is a strong similarity between sediment carried by the Neales River and sediment in the Davenport Ranges 4-28
Figure 4.23: Non Negative Principal Components Analysis (NNPCA) results for ASTER Neales Fan image showing poor discrimination of surface features. It identifies areas of erosion in yellow; relict surfaces in green; and channels in purple

igure 4.24: Independent Component Analysis (ICA) Neales Fan ASTER Image showing classification of the urface of the Neales Fan into several units. One of these appears to be a palaeochannel that bifurcates down the length of the fan (pink and pale blue). The yellow/green member is associated with a plateau on the northern ide of the Neales River. Similar material is evident on the southern side
igure 4.25: Rowan's Ratios of RGB muscovite/illite/smectite, calcite, alunite/pyrophillite for the Neales Fan howing a strong signal response to muscovite minerals across the scene and an association with sand dunes nd alunite/pyrophillite4-32
igure 4.26: TIR Silica Index for ASTER Neales Fan showing a strong signal response for areas covered in gibbe lains4-34
igure 4.27: Mahalanobis Distance Classification Image for Neales Fan ASTER scene, showing areas classified to three main units: areas of gibber plains (dark blue), areas undergoing active erosion (cyan) and areas that ave been reworked by fluvial/shoreline processes (red)4-35
igure 4.28: Non Negative Principal Component Analysis (NNPCA) of ASTER results for Sunny Creek showing unefields to the south with significant inputs of sand from reworked Jurassic sediments4-37
igure 4.29: Independent Component Analysis ASTER Image for Sunny Creek scene showing a difference in unefield composition between eastern dunes in green and western dunes in pink. Algebuckina sandstone is isplayed clearly in red. Orange defines erosional depressions that grade into alluvial depositional plains defined a pale green. 4-38
igure 4.30: Rowan's Ratios for RGB muscovite/illite/smectite, calcite, alunite/pyrophyllite for Sunny Creek STER scene, showing that the Davenport Ranges have a strong response in the calcite band ratio in green orresponding to the Skillogalee Dolomite Formation. This signal is also evident in the alluvial channels flowing om the ranges and is interpreted as the provenance for the sediment in these streams. Dunes in the east are ominated by alunite/pyrophillite signal in blue. Dunes in the west show a greater influence from huscovite/illite/smectite
igure 4.31: Rowan's Ratios for RGB muscovite/illite/smectite, dolomite, alunite/pyrophyllite for Sunny Creek STER scene showing a strong response from the alluvial fan sediments4-40
igure 4.32: Rowan's Ratios for RGB muscovite/illite/smectite, dolomite, calcite for Sunny Creek ASTER scene howing a mixed calcite and dolomite signal across the landscape. Alluvial fans have a strong muscovite mineral ignature
igure 4.33: Mahalanobis Distance Classification image for Sunny Creek ASTER scene, showing strong signal eparation between defined units. Proterozoic sediments of the Davenport Ranges are well-defined in blue as are algebuckina Sandstone sediments in red. Green defines areas associated with erosion of Bulldog Shale articularly in the Hawker Springs area. The eastern and western dunefields are clearly visible in yellow and cyar espectively. Alluvial channels are evident in maroon. These are flanked by alluvial erosional depressions in cora alluvial depositional plains in chartreuse. These two features are interspersed downstream and demonstrate the heterogeneous nature of the erosion/deposition cycle in the region. Surficial lags composed of degraded uvial sediment are evident in magenta. These mantle the area between Sunny and George Creeks and the area round Mount Anna where remnants of silcrete-cemented fluvial sediment are preserved and topographically everted in the landscape.
igure 4.34: Location diagram for JERS-1 SAR data showing the position of the JERS-1 SAR tiles used for nalysis (Landsat7 ETM mosaic courtesy of Exxon-Mobil)4-45
igure 4.35: JERS-1 Synthetic Aperture Radar Scene A, showing strong reflections from two areas of bedrock and exposed faulted margins, moderate reflection from stream beds, sand dunes, mound springs, escarpments and reas with high surface roughness and poor reflection from clay. Original Data © NASDA, MITI (1995) distributed y ACRES
igure 4.36: JERS-1 Synthetic Aperture Radar Scene B, showing high reflection from some areas of the Lake yre playa surface, streams, sand dunes, mound springs and topographically inverted silcrete (indicated by rrows); and poor reflection from claypans. Original Data © NASDA, MITI (1995) distributed by ACRES 4-48
igure 4.37: Scene A meld image JERS-1 SAR/ASTER showing distinction between geologic units within the lavenport Ranges. Claypans and springs are evident as red areas and floodplains flanking the Neales River are learly defined in dark blue. The surficial lags between George and Sunny Creek in blue are distinct from the lluvial fans flanking the ranges in purple

Figure 4.38: Scene B meld image JERS-1 SAR/ASTER showing the surface of the Neales Fan. Areas of gibber plain are green; yellow tends to highlight eroded alluvial channels and sand dunes. Alluvial sand plains are visib in orange and are associated with interdunal pans in blue. A palaeochannel can be seen diverging from the shall bend in Sunny Creek. Coastal dunes and palaeo-shorelines can be distinguished near the modern shore of Lake Eyre. Topographically inverted silcretes are distinguished in purple and pink along the northern edge of Sunny Creek
Chapter 5: Tectonic Influences on Geomorphology Figure 5.1: Location map showing the names and positions of tributary streams in the Umbum Creek Catchment
Figure 5.2: Drainage near Lambing Creek showing internal, diverted and convergent drainage related to faults
Figure 5.3: Trellis-like drainage along Browns Creek demonstrating drainage control from sand dunes5-
Figure 5.4: Divergent drainage between Sunny and Umbum Creeks showing the bounding faults that form the structural block
Figure 5.5: Asymmetric drainage along Davenport Creek indicating probable tectonic titing along the axis of the valleys
Figure 5.6: Oblique aerial photograph showing inherited meandering stream segments preserved in Proterozoic bedrock along Hawker Creek, indicating drainage antecedence prior to uplift of the ranges5-
Figure 5.7: Categorised stream plan morphology showing the dominant association of anastomosing channels with alluvial fans
Figure 5.8: Longitudinal profiles for streams flanking the Davenport Ranges. All streams show a positive deviation in their headwaters along the rangefront, evidence of movement along the Mt. Margaret Fault. Shaded areas are anastomosing reaches associated with alluvial fans
Figure 5.9: Longitudinal profile of southern streams showing a positive deviation in the medial section of each profile. This is interpreted as evidence of uplift in this zone. Shaded areas represent anastomosing reaches
Figure 5.10: Longitudinal stream profile of the Neales River showing the location of anastomosing reaches as shaded areas. The downstream boundary of each of these reaches corresponds to a change in slope and a physical nick point
Figure 5.11: Calculated slope image for the 9 second DEM. Proterozoic inliers and topographically inverted silcrete outcrops are prominent. The North bank of the Neales River has a steep slope and is associated with the Lake Eyre Fault. The Davenport Ranges has an escarpment on its eastern perimeter that is defined by the Levi and Mt. Margaret Faults. Movement along these faults has probably caused channel incision in Hawker and Lev creeks.
Figure 5.12: Calculated aspect image for the 9 second DEM showing two sets of aspect trends: a northwesterly trend and a southwesterly trend. These trends are probably a result of movement along underlying faults 5-1
Figure 5.13: Delimitation of catchments within the study area showing the locations of catchment boundaries.5-1
Figure 5.14: Classification results for Strahler stream ordering showing the Neales River to be the largest stream carrying the most flow in the study area
Figure 5.15: Morphometric analysis of all basins, displaying log-normal distribution5-1
Figure 5.16: Morphometric analysis of the Neales River Sub-basin, displaying log-normal distribution5-1
Figure 5.17: Morphometric analysis of the Umbum Creek Sub-basin, displaying log-normal distribution 5-1
Figure 5.18: Morphometric analysis of the Douglas Creek Sub-basin, displaying log-normal distribution5-1
Figure 5.19: Hyspometric curves showing convex-down shapes associated with high erosion rates5-2
Figure 5.20: Elevation histogram showing clusters of peaks in elevation at 2.5 m, 40-50 m and 100-150 m that ar possibly associated with palaeo-land surfaces and may indicate tectonic events
Figure 5.21: Relative Stream Power (RSP) showing a general increase in stream power downstream but with the highest values occurring east of the Lake Eyre Fault
Figure 5.22: Sediment Transport Capacity (LS) demonstrating increases in LS east of the Lake Eyre Fault 5-2

Figure 5.23 : Modeled results for flooding at a) –10 m AHD, b) -3.5 m AHD, c) +5 m AHD and d) +10 m AHD, showing the calculated flooding levels over modern topography
Figure 5.24: Location diagram showing the positions of RTK-DGPS survey transects
Figure 5.25: Selected RTK-DGPS surveys of the Neales River showing distinct valley asymmetry. See Figure 5.24 for profile locations
Figure 5.26: Selected RTK-DGPS cross-profiles of Umbum Creek showing asymmetry and flattening of relief downstream. See Figure 5.24 for profile locations
Figure 5.27: RTK-DGPS survey down the axis of the Neales Fan showing a gently sloping irregular surface with terracing. See Figure 5.24 for profile locations
Figure 5.28: RTK-DGPS topographic survey across the width of the Neales Fan showing the asymmetry of the fan surface. See Figure 5.24 for profile locations
Figure 5.29: Schematic geological sections of the Neales Fan looking north, showing terracing in the surface of the fan possibly associated with tectonic adjustment
Figure 5.30: Schematic geological section of the Neales Fan looking west, showing the strong north to south asymmetry and dissection developed across the fan surface
Chapter 6: Geological Structure Mapping Figure 6.1: Location map showing the location of Section C at the Neales Cliff
Figure 6.2: Structural data showing folding in the Etadunna Formation plotted on an ASTER greyscale image. (from Croke <i>et al.</i> , 1998) Section C is present in the east of the image. Etadunna Formation bedding changes dip from east to west and small-scale folds gently plunge approximately northwest: See Figure 6.1 for location 6-3
Figure 6.3: Representative facies profile and TL chronology of Quaternary main-channel fluvial (Warmakidyaboo Beds), lacustrine (Ghost Yard Beds) and aeolian sediments from the large meander bend at Section C. Facies are laterally continuous around the entire curvature of the bend. The basal deposits of Warmakidyaboo Beds directly overlie silicified deposits of the mid-Tertiary Etadunna Formation (from Croke et al., 1998)6-4
Figure 6.4: Bedding in Etadunna Formation truncated by unconformable Warmakidyaboo Beds, indicating deformation between the end of the Miocene and the Late Pleistocene. Location VWS002 on Figure 6.2 (0677491E 6884436N)6-5
Figure 6.5: Folded white clay interpreted as top of Etadunna Fm. indicating post-Miocene deformation. Location VWS005 on Figure 6.2 (0678059E 6884324N)6-5
Figure 6.6: Equal Area projection plot of assumed bedding planes in the Etadunna Formation showing poles plotted that equate to open, gently plunging folds. The calculated axis of the fold plots at 1°-139°6-6
Figure 6.7: Thrusted bedding planes present at the top of the Etadunna Formation and associated fracture planes, indicating tectonic activity following the deposition of the Etadunna Formation. Location VWS001 on Figure 6.2 (0677897E 6884393N)6-8
Figure 6.8: Rose diagram of observed fracture planes showing a conjugate set consistent with southwest-northeast compression
Figure 6.9: Location diagram of the Neales Fan showing the location of an inferred basement fault and associated meanders in the Neales River and Umbum Creek. Note that the Neales Overflow follows the same trend as the fault
Figure 6.10: Total station survey of sediments exposed in the Neales River bank showing deformation of Miocene sediments (Etadunna Formation) and no detectable deformation of Pleistocene sediments (Ghost Yard Beds) 6-12
Figure 6.11: Landsat 5TM image showing sand dunes formed on either side of the Neales River that may have blocked the channel as a result of the onset of aridity in Central Australia
Figure 6.12: Thalweg profile of the Neales River showing the decrease in slope trend across the meander. Arrows indicate nick points in the fluvial profile cut during falls in base level
Figure 6.13: DEM wireframe image of landscape at x50 vertical exaggeration showing the locations of potential faults influencing flow along the Neales River meanders. The faults form a series of parallel, narrow blocks that

fracture associated with the lineament. Here it follows the fracture and switches to the adjacent en-echelon fault.
Figure 6.14: Conceptual sketch of structural control of streams by block tilting. The stream flows along the base of one fault block until it reaches a fracture associated with the lineament. The stream exploits this weakness and switches across to the adjacent fault.
Figure 6.15: Extent of Etadunna Formation showing the close association of significant meanders on the Neales River and Umbum Creek with the change in lithology
Figure 6.16: Model of lithological control on stream planform showing energy expended as incision over erodable substrate in contrast to energy expended as lateral erosion over indurated substrate6-17
Figure 6.17: DEM of the Umbum Creek meanders constructed from 10 m spaced DGPS traverses and showing the control of streams by faults (dotted lines) below the meanders. (Courtesy M. Reilly & A. Hill)6-19
Figure 6.18: Umbum Creek thalweg profile from RTK-DGPS derived DEM showing the consistent trend of the profile associated with the meanders
Figure 6.19: Location map for Transect GPR01 showing the location of faults (f) that cross the transect and disrupt the sub-strata. The first 250m of Transect GPR01 is shown in Figure 6.206-21
Figure 6.20: The first 250m of GPR Transect GPR01 showing the offset within Pleistocene and Eocene units across the formation boundary, indicating Late Plesitocene tectonic activity6-22
Chapter 7: Regolith Geology Figure 7.1: Approximate distribution of surficial lags in the Umbum Creek Catchment in four zones, showing the association with landforms. Zone 1: alluvial fans; Zone 2: silcrete outcrops; Zone 3: fluvial sediments; and, Zone 4: lacustrine influence
Figure 7.2: General classification of surficial lags in the Umbum Creek Catchment. A) Zone 1: angular to subangular quartzite cobble and pebble lag with minor quartz, silcrete, hematite and lithic fragments (scale=5 cm); B) Zone 2: subangular to rounded silcrete pebble lag with uncommon quartz and hematite (scale=10 cm); C) Zone 3: subangular to subrounded cobble and pebble lag of mixed lithologies including silcrete, quartz, lithic fragments and hematite (notebook for scale); and, D) Zone 4: subangular to rounded pebble lag of mixed lithologies including silcrete, quartz, lithic fragments and hematite (hammer for scale)
Figure 7.3: Location of the silcrete-rich Mirackina Palaeochannel (MPC) showing a trend towards the silcrete outcrops within the study area
Figure 7.4: The Davenport Ranges displaying the Mt. Margaret Surface (arrow), a planation surface formed during the Jurassic and subsequently uplifted. (0602495E 6859603N)
Figure 7.5: Topographic section across the base of the Davenport Ranges from Douglas Creek to Lambing Creek looking west, derived from the DEM 9-S, showing the irregular nature of terraces. The section location is shown in the oblique view diagram above
Figure 7.6: Outcrop of Algebuckina Sandstone showing a coarse conglomeratic layer in the middle of the picture weathering to contribute rounded quartz clasts to the modern system (-0616126E 6830409N)7-11
Figure 7.7: Surficial lag composed of predominantly rounded quartz pebbles sourced from Algebuckina Sandstone and Cadna-Owie Formation, indicating the reworking of ancient fluvial sediments into the modern system (0616440E 6830635N)7-11
Figure 7.8: Tilted silcrete outcrops adjacent to the Davenport Ranges indicating tectonic activity following silcrete formation at a) Davenport Creek (view towards north 0616575E 6831738N) and b) Bulldog Creek (view towards north 0602476E 6861563N)
Figure 7.9: Detail of the Davenport Creek Silcrete showing tabular style at base, reed mould style in the middle and glaebular style at the top (hammer for scale ~0616575E 6831738N)7-14
Figure 7.10: Detail of Davenport Creek Silcrete showing glaebular style silcrete (hammer for scale ~0616575E 6831738N)
Figure 7.11: Detail of Davenport Creek Silcrete showing tilted reed mould style silcrete (hammer for scale ~0616575F 6831738N).

Figure 7-12: Silcrete cobble float (pale red clasts in picture) located atop mesas of highly ferrugenised Cadna-Owie Formation, indicating that the silcrete originated higher in the landscape than the hilltop (Scale 10 cm 0616139E 6830029N)
Figure 7.13: Silcrete pebble mantle overlying a pediment developed on Cadna-Owie Formation (foreground) and weathered Bulldog Shale (hills in background), indicating that the silcrete originated within the confines of the modern valley (View towards east 0619657E 6830240N)7-16
Figure 7.14: Rounded clast located atop Four Hills possibly indicating fluvial systems were active higher in the landscape. These outcrops were, however, used as stone tooling sites by Aboriginals and the rounded stones may represent anthropomorphically transported material for use as hammer stones (hammer for scale 0648859E 6846176N)
Figure 7.15: Comparison of spatial distribution of alluvial overbank RLU and the longitudinal stream profile of the Neales River showing the locations of major faults. The gradient across the Lake Eyre Fault shows a marked steepening and is associated with the termination of alluvial overbank sediments, indicating neotectonic control on landforms
Figure 7.16: Comparison of the spatial distribution of alluvial overbank RLUs and stream planform along Davenport and George Creeks, showing the locations of faults (f) associated with Brinkley Springs. The occurrence of alluvial overbank RLUs downstream of the faults indicates the development of landforms associated with neotectonic activity
Figure 7.17: ASTER/JERS-1 SAR meld image showing the neotectonic deviation of Sunny Creek. Movement along the fault has led to channel deviation in Sunny Creek and channel avulsion in Umbum Creek. This has left the palaeo-delta of Sunny Creek abandoned and it has since been degraded
Figure 7.18: Comparison of landform and topography showing the association of claypans (grey) with valleys parallel to the Neales River and Browns Creek, indicating the disruption of the land surface by faulting (f)7-25
Figure 7.19: Inferred palaeochannel across the surface of the Neales Fan defined by low thermal absorption on ASTER TIR. The low absorption areas are associated with claypans that represent a remnant channel plug
Figure 7.20: RTK-DGPS stream section Western Fan (indicated on Figure 5.24) across Umbum Creek channel showing the elevated eastern bank
Figure 7.21: Four Hills and surrounding landscape showing the dominance of silcrete surficial lags across the region and the significant topographic inversion of the silcrete caprocks (view towards east 0645725E 6845784N).
Figure 7.22: View towards the northwest from Four Hills displaying the low topography that dominates the region. Broad colluvial plains of red subangular to subrounded silcrete pebbles radiate from topographically inverted outcrops (car for scale 0645725E 6845784N)
Figure 7.23: Tilted 'reed mold' structures present in Pliocene silcretes of Four Hills, indicating tectonic activity in the Pleistocene following the formation of the silcretes in the Pliocene (hammer for scale 0646370E 6846176N)  7-29
Figure 7.24: Brecciation in silcrete at Four Hills (pen at top for scale 0645756E 6845838N)7-30
Figure 7.25: A silcrete boulder displaying clasts of silcrete matrix supported silcrete clasts. These clasts possess micro-clasts indicating at least three phases of silcrete generation (Scale 10 cm 0669537E 6845115N)7-30
Figure 7.26: In-situ silcrete outcrop showing water-worn surfaces with pebble-holes, indicating a river of significant power and volume flowed across this surface following the formation of the silcrete (scale 10 cm 0664347E 6846169N)
Figure 7.27: Highly polished, well-rounded pebbles located atop low hills of silcrete east of Four Hills. The polishing indicates these pebbles were transported in a high energy fluvial environment (scale 10 cm 0657021E 6845979N)
Figure 7.28: Backfilled burrow similar to <i>Muensteria isp.</i> probably created by dung beetle larvae and represents a fluvial or floodplain environment (scale 10 cm 0673836E 6846500N)
Figure 7.29: Location diagram of northeast trending longitudinal dunes shown in red
Figure 7.30: Location diagram of north-south trending dunes on the Neales Fan shown in red7-35

Figure 7.31: Idealised plan and cross-section of north-south trending dunes on the Neales Fan showing the 'buckshot' granule lags being reworked from the indurated core and redeposited downslope contributing to the gibber plain
Figure 7.32: Location for structurally related dunes on the Neales Fan shown in red7-37
Figure 7.33: Location for sand sheets shown in red
Figure 7.34: Location for coastal dunes
Figure 7.35: Modern drainage on the Neales Fan surface. Red marks drainage from east to west indicating a drainage divide (dashed line) formed across the mid-section of the fan7-40
Figure 7.36: Location of identified palaeochannel mounds (in black) on the Neales Fan, showing the association with coastal dunes that are interpreted to have been sourced from the palaeo-delta and meanders in the palaeochannel, possibly related to lithological control from the underlying Etadunna Formation (extent of Etadunna Formation marked by dotted line). For designated RLUs see accompanying Regolith-Landform map (Appendix C)
Figure 7.37: A: Log through an inverted palaeochannel on the Neales Fan. B: Landsat ETM7 image showing the surface expression of the outcrop. C. Interpretation of Landsat ETM7 image showing a meander of a palaeo-river on the Neales Fan (after Krapf and Lang, 2005)
Figure 7.38: Diagram of the Mirackina Palaeochannel (MPC) showing the location of remnant silcrete outcrops in the area and the proposed missing link between the silcretes of the MPC, Mt. Anna, Four Hills and Flint Mound 7-45
Figure 7.39: Carbonate-cemented fluvial conglomerate overlying Bulldog Shale, located in the headwaters of Sunny Creek, indicating the presence of Plio-Pleistocene palaeochannels through this area (foot for scale 0622618E 6830143N)
Figure 7.40: Colluvial mantle overlying carbonate-cemented fluvial sandstone overlying Bulldog Shale, located between Sunny Creek and Davenport Creek, indicating the presence of Plio-Pleistocene palaeochannels (looking east 0625233E 6830453N)
Figure 7.41: Colluvial mantle unconformably overlying fluvial sands and carbonate-cemented fluvial conglomerates overlying Bulldog Shale (notebook for scale, looking west 0625233E 6830453N)7-46
Figure 7.42: Comparison of gravimetric data to sand dune distribution showing the concentration of sand dunes (shown in yellow) over gravity lows and the general absence of sand dunes over gravity highs. Purple to blue colours show low gravity, yellow to red colours indicate high gravity7-47
Figure 7.43: Stream profile of Douglas Creek showing the positive deviation in the medial section of the stream associated with the gravitational high in red
Figure 7.44: Comparison of Total Magnetic Intensity and the distribution of gypsum patterned ground surfaces showing the general correlation between areas of low magnetic intensity (blue-green) and gypsum patterned ground surfaces. Areas of high magnetic intensity are recorded in orange-red7-50
Figure 7.45: Comparison of interpreted magnetic lineaments and regolith-landform units showing many instances of correlation between basement structure and surface landforms. Red lines indicate northwest-southeast trending magnetic lineaments; black lines indicate northeast-southwest trending magnetic lineaments. For designated RLUs see accompanying Regolith-Landform map (Appendix C)7-51
Chapter 8: Landscape Processes Figure 8.1: Diagrammatic representation of the overall process of sedimentary pyrite formation (from Berner, 1984)
Figure 8.2: Yardang near the western shore of Lake Eyre approximately 3 m in height (~0693996E 6852960N)
Figure 8.3: A: Classical model of deflation where erosion from wind and water cause progressive landscape lowering and the concentration of particles. B: Vegetation can prevent deflation by reducing wind speed below threshold velocities
Figure 8.4: Rainfall infiltration can cause cementing of surface fines and form a carapace, preventing wind deflation.

Figure 8.5: Physico-chemical weathering causing outcrops of in situ rocks to break down. Chemical weathering alters the parent material via oxidation/reduction and physical weathering via frost or salt expansion causes splitting as a result of volumetric change. The resultant angular fragments are distributed across the landscape by alluvial and colluvial processes.
Figure 8.6: Banded vegetation formation. Shallow unchannelised flow travels across the bare unvegetated run-of zones and is intercepted by the vegetated run-on zones8-8
Figure 8.7: View of banded vegetation developed on Cadna-Owie Formation as seen from the air, showing the horizontal development of run-off and run-on zones (~612800E 6845000N)8-9
Figure 8.8: Shrink-swell soil behaviour. Rainfall infitrates the soil, wetting clay particles and causing them to expand. This opens up small fissures around clasts and when drying of the soil causes expansion fines fall into these fissures gradually pushing the clasts upwards
Figure 8.9: Accretionary mantle formation. Wind deposits fine particles that work their way beneath surface stones. These aeolian deposits grow and form a pediment, lifting the gibber plain8-11
Figure 8.10: Pit location diagram showing the relationship between pit sites and surrounding landforms 8-13
Figure 8.11: Sedimentary profile of Pits JH002a & JH002. JH002a (left) was sampled through alluvial sand and is demonstrated to overlie the colluvial plain. JH002 (right) is composed of slightly weathered alluvial sediment with a gypsum-cemented layer. Landform and pit reference diagram (bottom) showing the location of pits in relation to landforms (see Figure 8.10 for pit locations)
Figure 8.12: Section location and logs for pits JH005, JH006 and JH007 showing the variability in substrate beneath a continuous style of gibber plain surface. JH005 has carbonate coated clasts, JH006 shows gypsum cementation and manganese nodule development and JH007 shows gypsum cementation, manganese nodule development and ripple cross-laminations (see Figure 8.10 for pit locations)8-15
Figure 8.13: Pit logs for JH008 and JH009 showing the variability in substrate beneath a similar gibber plain surface. JH008 shows a ripple cross-laminated sand with gypsum cementation and manganese nodule development while JH009 shows a pebbly sand overlying a green-grey clay (see Figure 8.10 for pit locations)
Figure 8.14: Pit logs for JH010, JH011 and JH012 showing the variability in the substrate between logs JH010 and JH011 that are beneath a similar gibber surface and the similarity between logs JH011 and JH012 that are beneath different gibber surfaces (see Figure 8.10 for pit locations)
Figure 8.15: Pie chart plot of sieve results for the soil sample from pit JH005 showing nearly a third of the sample composed of granule to pebble-sized clasts
Figure 8.16: Diagram of (A) a 1500 mL cube showing the length of each side and (B) the depth of material removed to create a 2 cm thick layer of gibber plain
Figure 8.17:Styles of gypsum sedimentation within the Umbum Creek Catchment.  A) Sheets of selenite deposited as veins within gypsum-rich shales (0647491E 6855831N).  B) Mesas formed by headward erosion of gypsum-cemented upland surface (~063557E 6881581N);  C) Gypsum-cemented sands (0641071E 6855015N);  D) Gypsum deposited as 'desert rose' crystals (0641071E 6855015N);  E) Gypsum-cemented rhizomorphs (~0678429E 6884775N);  F) Incipient gypsum precipitation in the soil profile (0659525E 6872501N);  G) Gypsum patterned ground (0635696E 6872482N); and,  H) Karstic dissolution of gypsum patterned ground (~0680062E 6872808N)
Figure 8.18: Morphological features of gypsum patterned ground.  A) Margin of a gypsite tableland displaying headward erosion (~637466E 6882051N);  B) Broad gypsum patterned ground tableland with preserved surface elevated above modern base level (~636227E 6872579N);  C) Gypsum patterned ground forming a mantle on Palaeoproterozoic quartzite inselberg (0634839E 6870580N);  D) Gypsum patterned ground forming a mantle on outcrop of Miocene silcrete (0664347E 6846169N);  E) Gypsum patterned ground surface showing meso-scale reprecipitation of gypsum downslope (0635696E 6872482N); and,  F) Gypsum patterned ground displaying mantling of the land surface. The mantle dips in both directions over undeformed sediment (0620755E 6855839N)

Figure 8.19: Processes of gypsum formation (from Aref 2003). (A) Rainfall causes leaching of gypsiferous shale that reprecipitates as a result of evaporation and the per ascensum mechanism. (B) Further rainfall leaches the gypsum crust reprecipitating gypsum as a subsurface profile via the per descensum mechanism. (C) Deflation erodes unconsolidated sediment producing gypsum-capped mesas. Runoff dissolves the gypsum crust that then reprecipitates downslope. Other mechanisms involve rainfall leaching of gypsum into the watertable and growth of gypsum below (D) the watertable forming a gypsum hardpan. (E) Deflation and erosion then result in topographic inversion of the gypsum hardpan.
Figure 8.20: Schematic diagram of erosion cell formation showing the development of production-transfer-sink zones and the interlinking of erosion cells in a mosaic (from Bourke and Pickup, 1999)
Figure 8.21: Erosion cell on the Neales Fan showing the demarcation between the production zone, the transfer zone and the sink with the direction of sediment transfer indicated
Chapter 9: Landscape Evolution Model Figure 9.1: Diagram of the development of the Umbum Creek Catchment with an oblique view to the north, showing the formation of alluvial fans along the rangefront following the uplift of the Davenport Ranges along the Mt. Margaret and Levi faults. The Neales River flows along the Lake Eyre Fault and diverges across the Neales Fan. A possible palaeochannel is present along the southern section of the study area9-6
Figure 9.2: Diagram of the development of the Umbum Creek Catchment with an oblique view to the north, showing the development of a gypsiferous horizon related to the topography9-6
Figure 9.3: Diagram of the development of the Umbum Creek Catchment with an oblique view to the north, showing renewed tectonism causing reactivation of the rangefront faults and alluvial fans, disruption of the land surface, stream diversions and mound spring development along faults and basement highs9-7
Figure 9.4: Diagram of the development of the Umbum Creek Catchment with an oblique view to the north, showing widespread erosion and dissection of the landscape. Silcrete outcrops become topographically inverted and colluvial processes become very active creating gibber plains and reworked Jurassic surface lags and stream inputs. Uplift along Douglas Creek causes doming of the land surface and concurrent incision of Douglas Creek.
Figure 9.5: Diagram showing the development of the Neales Fan during the Pliocene with an oblique view to the west, showing the trend of the Neales River structurally controlled by the Lake Eyre Fault and Browns Creek Fault9-10
Figure 9.6: Diagram showing the development of the Neales Fan in the Middle to Late Pleistocene with an oblique view to the west, showing renewed tectonism causing rearrangement of the landscape topography and tectonic cut-off of the Neales River, redirecting flow down the Umbum Overflow. Small-scale streams develop along faults
Figure 9.7: Diagram showing the development of the Neales Fan in the Late Pleistocene with an oblique view to the west showing abandoned channels forming the Neales Fan following stream capture of the Neales River via headward erosion from a smaller northern stream.
Figure 9.8: Diagram showing the development of the Neales Fan in the Holocene with an oblique view to the west, showing the growth of arid landforms under the influence of deflation. Sand is blown out of the palaeochannel leaving arcuate cobble covered mounds and forming sand dunes and sand sheets

# **LIST OF TABLES**

Chapter 1: Introduction  Table 1.1: Stratigraphic correlation chart for Tertiary units from various Australian sedimentary bas  & Pledge, 2000)	
Table 1.2: Quaternary climate change in Australia (from Williams, 1984)	1-11
Table 1.3: Episodes of Quaternary activity, western Lake Eyre (from Croke et al., 1996)	1-13
Table 1.4: Geomorphological summary of the Lake Eyre region based on work by various authors Twidale, 1967; Jessup & Norris, 1971; Croke <i>et al.</i> , 1996; Simon-Coinçon <i>et al.</i> , 1996; Alley, 1998; 2000.)	; and, Krieg,
Chapter 3: Seismicity in the Lake Eyre Region Table 3.1: Statistical Results (refer to Chapter 2.1 for statistical methods)	3-3
Table 3.2: Statistical results for the six year time-shifted frequency plots	3-8
Table 3.3: Statistical results for Inflow and Lake Level compared to Jackboot Bore RSWL	3-10
Table 3.4: Statistical results for comparison of regional datasets with Jackboot Bore RSWL	3-11
Chapter 4: Remote Sensing Table 4.1: Training Areas for Hawker Springs	4-28
Table 4.2: Training areas for Neales Fan	4-35
Table 4.3: Training areas for Sunny Creek	4-43
Chapter 5: Tectonic Influences on Geomorphology	
Table 5.1: Summary of Morphometric Analyses.	5-12
Table 5.2: Calculated values of rangefront sinuosity indicating that the rangefront is a young landso	
Table 5.3: Identified palaeo-hydrologic lake stages after De Vogel et al. (2004)	5-23
Chapter 6: Geological Structure Mapping Table 6.1: Trend and plunge of folds at the Neales Cliff	6-1
Chapter 8: Landscape Processes Table 8.1: Seive results for soil sampled from pit JH005	8-20
Chapter 9: Landscape Evolution Model Table 9.1: Landscape evolution summary chart for the Plio-Pleistocene. Times in bold represent d	
Table 9.2: Landscape evolution summary chart from the Permian to the Pliocene	9-13
Table 9.3: Main parameters indicating neotectonic activity	9-14
Table 9.4: Main factors affecting landscape evolution.	9-15

#### **ABSTRACT**

Landscape evolution is important for mineral and petroleum exploration concepts, especially in dryland continental settings. This study seeks to understand the main issues and controls on landscape evolution that have produced the regolith and young sediments around the western side of Lake Eye, in the arid heart of Australia.

Several methods were employed including satellite image analysis, geomorphometry, geological mapping, regolith mapping and surveying.

Outcomes indicate that the underlying structural fabric of the basement has controlled the development of the surface morphology of the Umbum Creek Catchment. The arrangement of basement faults is reflected in the distribution of surface landforms and in the topography of the land surface. Significant deformation of the Etadunna and Eyre formations indicate tectonic activity occurred at the end of the Miocene and was probably related to movement in the Lake Eyre Fault Zone. Pleistocene faulting is expressed as minor blind faulting associated with pre-existing basement faults. These faults remain active and current seismic activity is driven by changes in hydrostatic pressure (hydroseismicity). The scale of Pleistocene faulting and modern seismic activity demonstrates that since the Pliocene tectonic activity has been subdued.

Climate change caused landforms developed under wet conditions during the Palaeogene and Neogene to be preserved by the development of aridity in the Pleistocene. High erosion rates associated with tectonism and the onset of aridity in the Pleistocene led to topographic inversion of many features. Palaeo-Proterozoic inliers formed inselbergs, silcrete outcrops formed capstones, gypsum hardpans protected underlying sediment from erosion creating plateaux of gypsum patterned ground and palaeo-channels on the Neales Fan were eroded to make heavily armoured mounds and associated sand dunes and sand sheets.

The dominant factor influencing the evolution of the landscape in the Umbum Creek Catchment was the deposition of sedimentary sulphides within the Bulldog Shale. The excess sulphur that this sediment supplied to the landscape over time created the necessary conditions for the formation of a range of landscape features that would not otherwise exist. Weathering, oxidation and leaching of the sedimentary sulphides led to the development of silcrete. Subsequent weathering and tectonic activity led to the breakdown of the silcrete and the distribution of silcrete pebbles widely across the landscape forming gibber plains. Sulphur from the Bulldog Shale continued to contribute to the landscape forming intra-formational gypsum and precipitating as gypsum hardpans.

This study has implications for petroleum exploration in dryland continental settings as potential reservoirs may be affected by secondary diagenetic processes, such as the formation of gypsum or silcretes, that could act as baffles or result in reduced porosity within the reservoir.

The broad-scale architecture of fluvial systems, like the Neales Fan, may not conform to traditional fan-shaped models being, instead, comprised of structurally rearranged channels.

In terms of earthquake risk assessment, the identification of hydroseismicity active within the Lake Eye Basin allows for a new level of predictability of earthquake behaviour within Central Australia.

#### **DECLARATION**

This work contains no material which has been accepted for the award of any other degree or diploma in any university or other tertiary institution and, to the best of my knowledge and belief, contains no material previously published or written by another person, except where due reference has been made in the text.

I give consent to this copy of my thesis being made available in the University Library.

The author acknowledges that copyright of published works contained within this thesis (as listed below) resides with the copyright holder/s of those works.

#### **ACKNOWLEDGMENTS**

I would like to acknowledge the many people who assisted me during this research. A big thank you goes to my fellow field workers and field hands, Mark Reilly, Saju Menacherry, Ian Moffat, Adam Furbank and Tamara Baillie, for their time and effort in the northwestern deserts. Special thanks goes to my supervisors for their encouragement and support throughout this project: Simon Lang, Steve Hill, Tobi Payenberg, Carmen Krapf and Pat James.

The enthusiasm and feedback from international and local researchers was invaluable in the conduct of this work and the help of John Holbrook, John Fisher, Stephen Hasiotis, Travis Gotch, Megan Lewis and Jochen Kassan is gratefully acknowledged.

The Cooperative Research Centre for Landscape, Environments and Mineral Exploration and the Lake Eyre Basin Analogue Research Group Phase II Consortium are acknowledged for providing the funding to conduct this research.

Kidman & Co. is acknowledged for providing access to their stations and for their help and assistance while conducting fieldwork.

A final special thanks goes to the residents of William Creek for their support throughout my PhD. Without the aid of the good people of William Creek this project would not have been possible.

This thesis was professionally edited for spelling and grammar by Carolyn Lang of Geosed Consulting and Editing Services